

Cloud base signature in transmission spectra of exoplanet atmospheres

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Abstract

We present an analytical model for the transmission spectrum of a transiting exoplanet, showing that a cloud base can produce an observable inflection point in the spectrum. The wavelength and magnitude of the inflection can be used to break the degeneracy between the atmospheric pressure and the abundance of the main cloud material, however the abundance still depends on cloud particle size. An observed inflection also provides a specific point on the atmospheric P-T profile, giving us a “thermometer” to directly validate or rule out postulated cloud species. We will apply the model to the transit spectrum of HD189733b.

1 Introduction

Transmission spectra of transiting exoplanets can provide a constraint on the position and type of cloud layers, temperature profiles, and ultimately, the abundance of cloud-forming species. These properties are quite uncertain (Madhusudhan and Seager 2010, 2011; Moses et al 2012; Marois et al 2008), but are fundamental to characterizing exoplanet atmospheres.

Perhaps the most convincing evidence of high altitude hazes on any extrasolar planet to date is found in the case of the transiting hot Jupiter HD 189733b, which is a 1.1-Jupiter mass planet orbiting a bright nearby K star and an excellent target for detailed atmospheric studies. This planet is notable because its transit radius follows a smooth wavelength dependence (Pont et al. 2008, 2013; Sing et al 2011), and lacks signatures of molecular or atomic absorption, although Na and K are detected at high spectral resolution (Pont et al. 2013, Huitson et al. 2012). A currently popular explanation for the HD 189733b transmission spectrum is a haze of small particles, extending five atmospheric scale heights (Pont et al. 2013), that are better scatterers than absorbers (Lecavelier des Etangs et al.

2008, henceforth L08); that is, the particle’s scattering efficiency is large compared to its absorption efficiency. This property, providing a λ^{-4} opacity dependence, requires a material with an extremely small imaginary index of refraction (to avoid dominance by absorption, with its $1/\lambda$ wavelength dependence; see section 2) and MgSiO_3 was suggested as a candidate. Other silicates, however, also condense in similar regions, as does Fe, so further exploration of this hypothesis seems appropriate. Here, we present a model for interpreting the transmission spectra of transiting planets, extending the models of Fortney (2005; henceforth F05) and L08 to incorporate condensation cloud bases.

2 Planet and cloud geometry

We consider an atmosphere containing a cloud formed from condensed particles at an altitude defined by the intersection of the T-P profile of the atmosphere and the condensation vapor pressure of the material. The typical vertical structure of a cloud is a sharp base at the condensation temperature and pressure (T,P), with a more gradual vertical falloff in density (Ackerman and Marley 2001, henceforth AM01). For simplicity we have assumed a single condensate.

A photon that traverses this atmosphere along the line of sight in transmission geometry is affected by gas and cloud particle opacities. However, the opacity for photons with wavelengths that can reach altitudes underneath the cloud base is reduced as it crosses the region below the cloud base, where only gas opacity contributes. This gap in opacity between $-x_c$ and x_c (see Figure 1) shows up as an inflection point in the transmission spectrum, as described below.

The number density n_g of gas molecules with scale height H is a function of radial distance R from the center of the planet, as measured relative to some reference radius a :

$$n_g(R) = n_0 e^{-(R-a)/H} = n_0 e^{-z/H} \quad (1)$$

The cloud particle number density n_c is defined by a fraction ζ and for simplicity the particles are assumed to have the gas scale height (*i.e.*, are very small; AM01):

$$n_c(R) = \zeta n_g(R) = \zeta n_0 e^{-(R-a)/H} \quad (2)$$

For the case of MgSiO_3 (enstatite), we follow L08 in assuming it is limited by the number of magnesium atoms; thus

$$\zeta = 2\zeta_{Mg} \mu_{\text{MgSiO}_3} m_p / \mu_{\text{H}} \rho V, \quad (3)$$

$$\tau_h = \sigma_g(\lambda)n_0e^{-\frac{z}{H}}\sqrt{2\pi aH} + \sigma_c(\lambda)\zeta n_0e^{-\frac{z}{H}}\sqrt{2\pi aH} \left(1 - \operatorname{erf}\left(\frac{x_c}{\sqrt{2aH}}\right)\right), \text{ if } z < z_c, \quad (7)$$

where from geometry, $x_c = \sqrt{2a(z_c - z)}$ (Figure 1, assuming $(z, z') \ll a$). L08 calculated a nearly model-independent equivalent optical depth $\tau_{eq} \approx 1/2$ for a range of atmospheric scale heights, where the planet with its translucent atmosphere produces the same absorption depth as a sharp occulting disk. The altitude z at which $\tau_h(z, \lambda) = \tau_{eq}$ is then:

$$z = H \ln \left[\frac{P_0}{kT\tau_{eq}}\sigma_g(\lambda)\sqrt{2\pi aH} + \frac{\zeta P_0}{kT\tau_{eq}}\sigma_c(\lambda)\sqrt{2\pi aH} \left(1 - \operatorname{erf}\left(\frac{x_c}{\sqrt{2aH}}\right)\right) \right] \quad (8)$$

We can also rewrite equation 6 to express the pressure $P(z) = P_0e^{-z/H}$ of the hazy atmosphere above the cloud base and at $\tau = \tau_{eq}$ as:

$$P(z, \zeta, T, \sigma_c) = \frac{kT\tau_{eq}}{\sqrt{2\pi aH}(\sigma_g(\lambda) + \zeta\sigma_c(\lambda))} \quad (9)$$

The gas reference pressure P_0 is chosen at reference wavelength $\lambda_0 = 750\text{nm}$ (λ_0 is arbitrary and we choose this value to be consistent with L08) where the altitude $z = 0$ or $R = a$, and $P_0 = kT\tau_{eq}/\sigma_g(\lambda_0)\sqrt{2\pi aH}$.

For a single component atmosphere where either gas or particles dominate, equation 8 can be written (neglecting constants) as $z \sim H \ln(\lambda^{-\alpha}) \sim -H\alpha \ln(\lambda)$, where $\sigma(\lambda) \sim \lambda^{-\alpha}$. Therefore for a given temperature, the transit curve is a log- λ function with slope ranging from flat for large particles ($\alpha = 0$), to a steeper value for gas molecules or non-absorbing Rayleigh scatterers with constant real refractive index ($\alpha = 4$). The scale height H also contributes to the slope (L08).

To the extent that the particle refractive indices are wavelength dependent, more detailed variations can occur in the transit spectrum. For instance, small particles in the Rayleigh regime show different wavelength and particle size dependence if they are essentially lossless (*eg.* van de Hulst 1957):

$$\sigma_c(r, \lambda) \approx Q_s\pi r^2 = \frac{24\pi^3(n_r^2 - 1)^2}{\lambda^4(n_r^2 + 2)^2}V^2 \quad (10)$$

or if they are even slightly absorbing:

$$\sigma_c(r, \lambda) \approx Q_a\pi r^2 = \frac{36\pi n_r n_i}{\lambda(n_r^2 + 2)^2}V. \quad (11)$$

Meanwhile for the gas,

$$\sigma_g(\lambda) = \sigma_0 \left(\frac{\lambda}{\lambda_0} \right)^{-4}, \text{ where } \lambda_0 = 2.52 \times 10^{-28} \text{cm}^2 \text{ and } \lambda_0 = 750 \text{nm} \quad (12)$$

3 Model Results

3.1 Inflection points

We illustrate the observable effect of a cloud base with a simple toy model calculation using equation 8, with a constant temperature atmosphere consisting of Rayleigh scattering H_2 molecules, and MgSiO_3 grains for the cloud constituent. In Figure 2 (left), we show the altitude (related to apparent radius) at which $\tau_h = \tau_{eq}$. Results are shown for cloud bases at two different pressures; note the inflection points which occur at two different corresponding wavelengths λ_D . For lines of sight passing above the cloud base (altitudes $z > z_c$, sensed by wavelengths $< \lambda_D$), both gas and cloud particles contribute to the opacity, leading to more extinction at shorter wavelengths than the gas-only curve (dotted blue). On lines of sight with altitudes below the cloud base, sensed by wavelengths $> \lambda_D$, primarily gas opacity contributes and the red curve falls onto the gas curve. In section 4 we will show the effect of actual T-P profiles.

A cloud base inflection point, if observed, has interesting implications about cloud particle properties and atmospheric properties in general. To illustrate this, we rewrite equation 8 in simpler form:

$$z = H \ln(A + B) - \frac{x_c}{\sqrt{2aH}} \quad (13)$$

$$A = \frac{\sigma_g(\lambda)P_0\sqrt{2\pi aH}}{kT\tau_{eq}}, \quad B = \frac{\sigma_c(\lambda)\zeta P_0\sqrt{2\pi aH}}{kT\tau_{eq}}, \quad (14)$$

$$\frac{B}{A} = \frac{\sigma_c(\lambda)\zeta}{\sigma_g(\lambda)}, \text{ and } D = H \ln\left(1 + \frac{B}{A}\right) \quad (15)$$

where D is the height of the inflection step. In Figure 2 (left panel) the transit altitude profile is broken up into two regions: the spectral region where gas and cloud particle opacities contribute (green curve at $\lambda < \lambda_D$) is approximated by $H \ln(A + B)$ and the region sampling beneath the cloud base, where gas opacity dominates, by $H \ln(A)$. The drop at λ_D from $H \ln(A + B)$ to $H \ln(A)$, corresponding to detection of the cloud base, is the inflection height D . Shown in Figure 2 (right panel) is the shift of the inflection point in wavelength depending on particle abundance ζ while keeping the cloud base pressure constant. Higher ζ results in

larger opacity, so longer λ is needed to reach the altitude/pressure where the cloud base is located; this then also results in a larger value of D (equation 15).

The two critical properties - the drop D and the wavelength λ_D - allow us to determine the cloud material abundance (a function of particle size) and cloud base pressure. The abundance ζ , P_0 , and σ_c are degenerate in the term B above, but if we can detect a cloud base in the transit spectrum, have a candidate condensate and T-P curve, and assume a particle radius r to calculate σ_c , then the degeneracy can be broken and limits can be placed on the abundance ζ for the candidate material as shown in section 4.

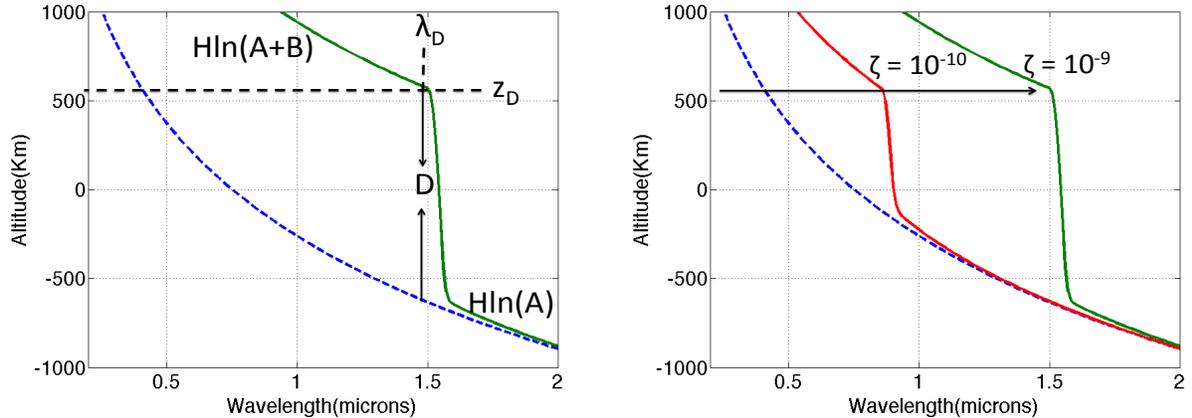


Figure 2: (left) The transit altitude profile showing transition from the spectral region where cloud particle opacity dominates ($z = H \ln(A+B)$) to the region sampling beneath the cloud base, where gas opacity dominates ($z = H \ln(A)$). The inflection point defines two critical parameters - the wavelength λ_D and the drop D (see section 3). The value of λ_D depends on the pressure at the cloud base. (Right) The value of λ_D also depends on particle abundance ζ , and a higher ζ also results in a larger drop D .

3.2 Cloud base and cloud top

Figure 3 shows two candidate T-P profiles for HD189733b and several arbitrary profiles. Even the two actual candidates are very different, indicative of the current uncertainties. We also show vapor pressure curves for several different materials; a cloud base forms when the decreasing temperature with altitude crosses a vapor pressure curve. Moreover, we could have a cloud with well-defined base *and* top, as in the case of the MgSiO_3 condensation line crossing the blue T-P profile with inversion, at two places (Figure 3, T-P profile is from Pont et al. (2013) using a model from Heng et al. (2012)). For simplicity we can break up the calculation for this scenario into three opacity regions: as wavelength increases, photons traverse deeper into the atmosphere. They first encounter purely gas opacity, then below

the cloud top gas and condensate opacities both contribute, and finally below the cloud base gas opacity dominates again.

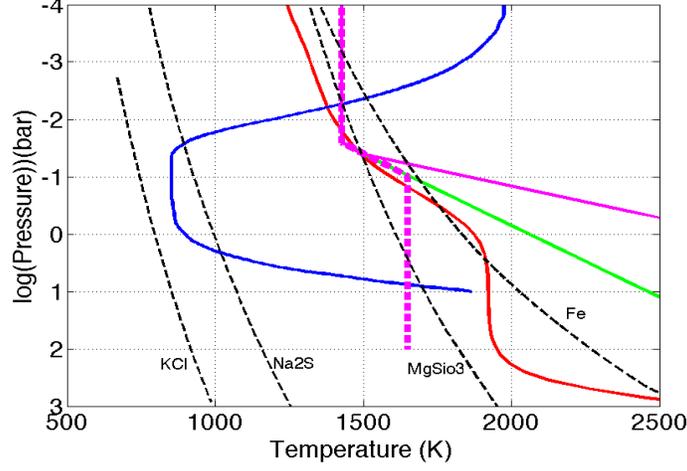


Figure 3: Alternate T-P profiles for HD189733b: the red curve is shifted by +250K from Fortney et al. (2008; see section 4) and the blue curve is from Pont et al. (2013). The green and magenta curves are artificial profiles we constructed to study the effect of the temperature gradient on the spectrum. We also show condensation curves (dashed black) for potential cloud forming candidate materials; cloud bases form when T-P curves intersect these vapor pressure curves.

Region 1, $z > z_t$:

$$z \sim H \ln \frac{P_0}{kT\tau_{eq}} \sigma_g(\lambda) \sqrt{2\pi aH} \quad (16)$$

Region 2, $z_c < z < z_t$:

$$z \sim H \ln \frac{P_0}{kT\tau_{eq}} \sigma_g(\lambda) \sqrt{2\pi aH} + \frac{\zeta P_0}{kT\tau_{eq}} \sigma_c(\lambda) \sqrt{2\pi aH} \operatorname{erf} \frac{x_t}{\sqrt{2aH}} \quad (17)$$

Region 3, $z < z_c$:

$$z \sim H \ln \frac{P_0}{kT\tau_{eq}} \sigma_g(\lambda) \sqrt{2\pi aH} + \frac{\zeta P_0}{kT\tau_{eq}} \sigma_c(\lambda) \sqrt{2\pi aH} \operatorname{erf} \frac{x_t}{\sqrt{2aH}} - \operatorname{erf} \frac{x_c}{\sqrt{2aH}} \quad (18)$$

4 Model comparisons with data

Recently, Pont et al. (2013) published a combined transit spectrum of HD189733b from their newly tabulated data from HST, ranging from UV to Infrared. This data set covers a larger range of wavelengths, at better resolution than previous observations, and may show hints of

inflection points. The spectrum can be described with a combination of two different slope $\ln(\lambda)$ curves as shown in Figure 4, where the slope change near $0.6\mu\text{m}$ may also be marked by an apparent drop similar to our inflection height D . However, the temperature needed to match the slope at short wavelengths is 2000K (Pont et al 2013), too hot for MgSiO_3 to condense. We return to this point below.

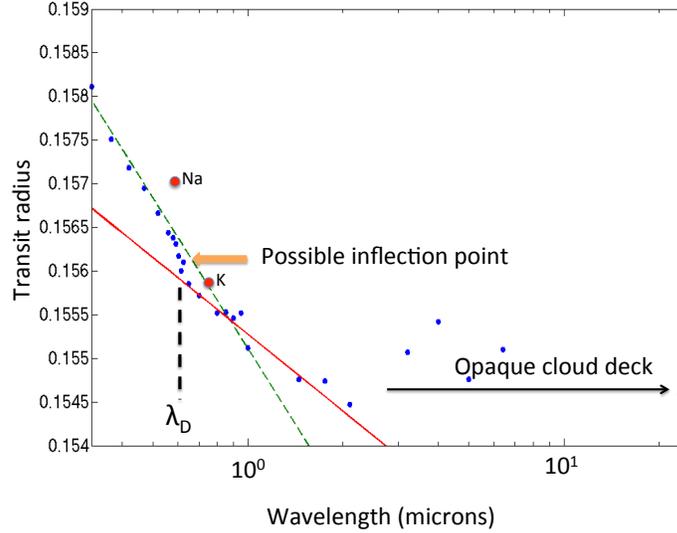


Figure 4: HD189733b data (taken directly from table 6 of Pont et al. 2013, black) compared with arbitrary log functions (red and green) with two different slopes (x -axis is shown with log scale). Na and K gas absorption lines are shown as red dots. Some have suggested that the data at $\lambda > 1\mu\text{m}$ suggest an opaque cloud deck.

4.1 Summary of procedure

Before comparing the observations with model transmission curves we summarize the procedure as follows:

1) Calculate cloud base temperature T_c from the transit slope (T_c needs to be consistent with the presumed cloud material).

$$T_c = \frac{\mu g}{k} \frac{dz}{d\lambda} \left(\frac{d \ln(\sigma_g + \zeta \sigma_c)}{d\lambda} \right)^{-1}. \quad (19)$$

2) Using equation 15 with the observed inflection point wavelength λ_D and drop D , calculate $\zeta \sigma_c(\lambda_D)$. Since we don't know the size of the cloud particles (beyond the hypothesis that $r \ll \lambda$), the particle abundance and particle cross section remain degenerate:

$$\zeta \sigma_c = \sigma_g(\lambda_D)(e^{D/H} - 1). \quad (20)$$

3) We then rewrite equation 3, using the product $\zeta\sigma_c$ from above, and substitute equation 10 (the nominal case where the particles are Rayleigh *scatterers*), to determine the chemical abundance of magnesium ζ_{Mg} in the cloud:

$$\zeta_{Mg} = \frac{\rho\zeta\sigma_c}{200m_p} \frac{V}{\sigma_c} = \frac{1\text{cm} (n_r^2 + 2)^2}{64\pi^4 (n_r^2 - 1)^2} \frac{\rho\zeta\sigma_c}{m_p} \frac{\lambda_D}{1\mu\text{m}}^4 \frac{0.01\mu\text{m}}{r}^3 \quad (21)$$

4) Substituting λ_D , T_c and $\zeta\sigma_c(\lambda_D)$ into equation 9, with $\tau_{eq}=0.5$, we calculate the cloud base pressure P_c . Then we can calculate a value for the reference radius a which we can use to fit our model to data at other wavelengths (given as apparent radius). Specifically, $a = R_{\lambda_D} - z_c$ where $z_c = H \ln(P_0/P_c)$.

Using the observed slope for $\lambda < \lambda_D$, equation 19 gives $T_c = 2000\text{K}$, as in Pont et al (2013). Since this temperature is too hot to sustain an MgSiO_3 cloud, we adopt a lower temperature $T_c = 1400\text{K}$ where MgSiO_3 remains solid. This discrepancy might be explained by a combination of uncertainty in the observed slope or in the value of g , and neglect of higher order terms in equation 8 propagating into equation 19. The corresponding cloud base pressure from equation 9 is $P_c = 10^{-1.3}$ bar. The magnesium abundance ζ_{Mg} calculated from equation 21 is five orders of magnitude lower than cosmic abundance assuming haze particles with $r \sim 0.01\mu\text{m}$.

4.2 Specific comparisons

Finally, we used various T-P profiles shown in Figure 3 to create cloud models for comparison with the data of Pont et al (2013), as shown in Figure 5. The first calculation in Figure 5 (left) uses the red T-P profile from Figure 3, which was shifted by 250K from Fortney et al (2008) so that the cloud base pressure, set by the altitude where the MgSiO_3 condensation line crosses the T-P profile, is equal to our calculated P_c . We can see that the model inflection point coincides with the possible inflection point in the data, but the model slope at shorter wavelengths is too flat. We even used actual, wavelength-dependent refractive indices of MgSiO_3 to get the fit shown, but these didn't change the slope much from some typical constant value. Moreover, the model slope at $1 - 2\mu\text{m}$ wavelengths is too *steep*.

We created artificial temperature profiles to investigate the effect further. These artificial T-P profiles (Figure 5, center) also have their cloud base and top defined by the altitude/pressure where they cross the MgSiO_3 condensation line. The dotted magenta curve is the closest in spirit to the modified Fortney et al. (2008) T-P curve, but being colder, has a flatter slope at long wavelengths. However it is actually slightly cooler just above the cloud base, so fits less well at short wavelengths.

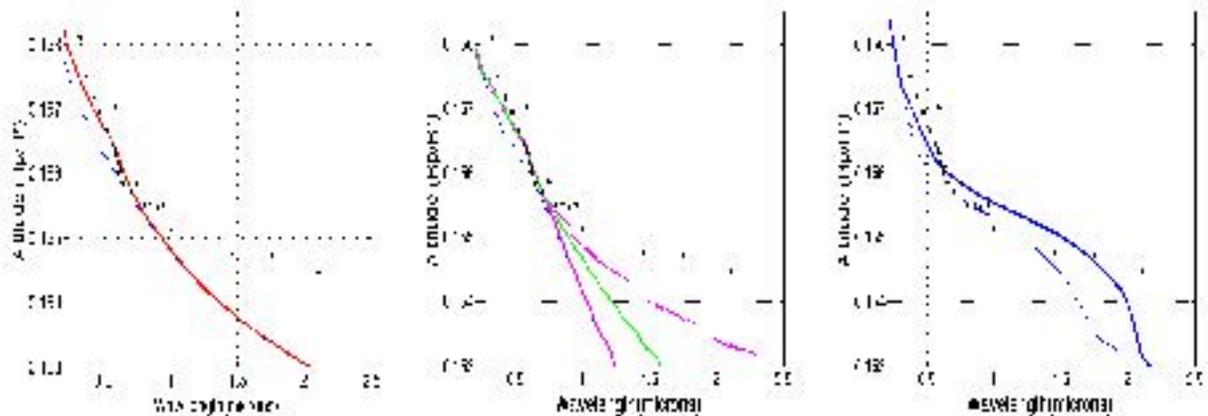


Figure 5: (left) Ratio of planet radius R_p to host star radius R^* vs. wavelength. Data from Pont et al (2013) (black symbols) are fitted with our cloud model calculation using the red T-P profile in Figure 3 and a cloud base pressure of $\sim 10^{-1.3}$ bar set by the altitude where the MgSiO_3 condensation line crosses the red T-P profile. The gas curve (dotted blue) is also plotted for reference. (Center) Pont et al. data is fitted with our cloud model calculation using the artificial profiles in Figure 3 with a cloud base and top pressures of $10^{-1.4}$ and $10^{-2.3}$ bar defined by the position where the MgSiO_3 condensation line crosses the T-P profiles. (Right) Pont et al. data is fitted with our cloud model calculation using the blue T-P profile in Figure 3 with cloud base and top pressures of $10^{0.8}$ and $10^{-2.3}$ bar set by where the MgSiO_3 condensation line crosses the blue T-P profile.

Figure 5 (right) shows the predictions of the blue T-P profile (Pont et al., 2013). This profile produces a cloud top inflection point at about $0.3\mu\text{m}$, and the steep slope at the shortest wavelengths is caused by hot H_2 gas, but the enstatite cloud *base* is too deep in the atmosphere, giving $\lambda_D > 2\mu\text{m}$ and a transit that would obscure the K gas absorption line.

Overall, the possible $0.6\mu\text{m}$ inflection may be consistent with an enstatite haze with abundance much smaller than cosmically available, for a T-P profile not too different from that of Fortney et al. 2008, but the steepness of the wavelength dependence at $\lambda < \lambda_D$, if interpreted as a temperature of 2000K, is not consistent with MgSiO_3 as the responsible haze forming material. The blue T-P profile has a thick enstatite cloud and a warm upper atmosphere, but is not consistent with a $0.6\mu\text{m}$ inflection and obscures the K line.

On a speculative note, we remark that the low abundance (relative to cosmic) we obtain for the tiny haze particles which cause the inflection at $\lambda_D \sim 0.6\mu\text{m}$ and dominate at shorter wavelengths, *and* the moderately flat spectrum for $\lambda > \lambda_D$, might suggest a vertically extended haze of tiny (perhaps enstatite) particles overlying a highly settled enstatite cloud with large particles containing most of the Mg mass, lying very close to z_c (like the settled clouds of AM01). The opacity of this settled layer might not be large, depending on the

particle size, but could be wavelength independent, with a small effective scale height, and might create the moderately flat spectrum for $\lambda > \lambda_D$. As the data improves, perhaps these possibilities may be explored.

5 Summary

We have extended the analyses of LeCavelier et al. (2008) to demonstrate the observable effects of condensation cloud bases on transit spectra. We showed how the signature of a condensation cloud base is an inflection or step at some λ_D in the apparent radius spectrum. The inferred T-P value at the inflection provides important constraints on condensate identification. The magnitude and wavelength of the inflection can provide an estimate of the abundance of the condensate (however, this value depends on the poorly known haze particle size). The single cloud base development presented here is easily extended to multiple cloud bases. The spectrum of HD189733b could be viewed as having one or more inflections; more data are needed.

References

- Ackerman, A. & Marley, M., 2001 APJ., 556, 872.
- Bohren, C. and D. R. Huffman, 1983, New York: Wiley.
- Fortney, J. J., 2005, M.N.R.A.S. 364, 649-653.
- Fortney, J.J., Lodders, M.S. Marley, R.S Freedman 2008, APJ, 678, 1419-1435.
- Heng K., Hayek W., Pont F., Sing D. K., 2012, MNRAS, 420, 20.
- Huitson, C. M.; Sing, D. K.; Vidal-Madjar, A.; Ballester, G. E.; Lecavelier des Etangs, A.; Dsert, J.-M.; Pont, F., 2012, MNRAS 422, 2477-2488.
- Lecavelier Des Etangs, A.; Pont, F.; Vidal-Madjar, A.; Sing, D., 2008, A&A 481, L83-L86.
- Madhusudhan, N. and S. Seager, 2010, APJ, 725, 261-274.
- Madhusudhan, N. and S. Seager, 2011, ApJ, 729, article id. 41.
- Marois, C., Macintosh, B., Barman, T., Zuckerman, B., Song, I., Patience, J., Lafrenire, D., & Doyon, R., 2008, Science, 322, 1348.
- Morley, C., Fortney, J. J., Kempton, E.; Visscher, C. W.; Marley, M. S. 2013, ApJ, 775, 33.
- Moses, J.I.; Richardson, M. R.; Madhusudhan, N.; Line, M. R.; Visscher, C.; Fortney, J. J., 2012; DPS meeting #44, #103.02.
- Pont, F.; Knutson, H.; Gilliland, R. L.; Moutou, C.; Charbonneau, D., 2008, MNRAS, 385, 109-118.

Pont, F., Sing, D. K., Gibson, N. P., Aigrain, S., Henry, G., Husnoo, N., 2013, MNRAS, 432, 2917-2944.

Sing, D. K., F. Pont, S. Aigrain, D. Charbonneau, J.-M. Desert, N. Gibson, R. Gilliland, W. Hayek, G. Henry, H. Knutson, A. Lecavelier des Etangs, T. Mazeh and A. Shporer, 2011, 416, 1443-1455.

Van de Hulst H. C. (1957) Wiley and Sons, New York.